Developing the representation of standing water, inundation and river routing in JULES

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This talk describes work in progress in two areas:

River routing and inundation

- Introducing the CaMa-Flood model into JULES
  - an alternative to existing parameteristions

Effects of inundation

- Effects on evaporation, infiltration, etc.
  - not currently represented

### CaMa-Flood

## Catchment-based Macro-scale Floodplain model

Yamazaki et al., 2011, A physically-based description of floodplain inundation dynamics in a global river routing model, Water Res. Res. Yamazaki et al., 2013, Improving computational efficiency in global river models by implementing the local inertial flow equation and a vector-based river network map, Water Res. Res.

## A 1-D model of river routing and inundation

Uses the local inertial equation – improved physics over kinematic or diffusive wave models Calculates water depth, then inundation estimated using sub-grid topography Uses an adaptive timestep approach on a grid or unit catchment basis.



#### **CaMa-Flood**

The local inertial equation allows backwater effects to be modelled. A downstream water level (e.g. sea level) can be used as a boundary condition – can capture effects of marine storm surge on rivers.

Channel bifurcation has been parameterised – as the water level rises some of the flow can access different flow paths.



CaMa-Flood is being implemented as an option in JULES – alongside TRIP and RFM.

## JULES-CaMa-Flood

Test results: comparing JULES-CaMaFlood with CaMa-Flood v3.96a. Both models are forced with runoff on a 0.25deg grid. 1 year run "from cold".



#### Effects of standing water

At present overbank inundation is purely diagnostic in JULES – the flood water does not affect any other aspect of the model.

In reality, standing water affects many aspects of hydrology and surfaceatmosphere fluxes.



## **Evaporation of standing water**

This is similar to the existing "lake evaporation" except that it depletes a finite store.

$$E = E_{\text{pond}} + E_{\text{can}} + E_{\text{soil}}$$

$$E = f_{\text{p}} \frac{\rho \, \delta q}{r_{\text{a}}} + (1 - f_{\text{p}}) f_{\text{c}} \frac{\rho \, \delta q}{r_{\text{a}}} + (1 - f_{\text{p}}) (1 - f_{\text{c}}) \frac{\rho \, \delta q}{r_{\text{a}} + r_{\text{c}}}$$
Fraction with  $E_{\text{can}}$   $f_{\text{c}} = \frac{c}{c_{\text{max}}}$ 

$$f_{\text{p}} = \frac{z_{\text{pond}}}{z_{\text{veg}}}$$

= depth of standing water/ height of vegetation

$$f_{\rm p} = \frac{z_{\rm pond}}{10z_{0,{\rm m}}}$$
 for non-veg tiles = depth/10\*roughness

We need a depth of standing water...

#### **Evaporation of standing water**

At present JULES cannot hold any water on the soil surface – any input that cannot infiltrate immediately forms surface runoff.

To allow testing I have simply diverted a constant fraction of throughfall into a new surface store.

The surface store can evaporate or infiltrate (at a fixed rate up to Ksat).

Or for testing...prescribe a constant depth of water.

Total evaporation from:



#### How to include standing water ?



Microtopography is assumed to be distributed normally. With assumptions about how patches are connected, we can calculate the extent of surface water and a threshold storage above which surface runoff occurs.

Needs to be made consistent with the rest of the model, including surface runoff (and rainfall distribution!), and groundwater and overbank inundation...

CaMa-Flood parameterisation of routing and overbank inundation

- Initial implementation working
- Needs tidying, testing, etc.

Representing the wider effects of inundation

- Standing water represented in surface fluxes
- Needs further work to introduce a store of standing water ...which needs to fit with with existing parameterisations