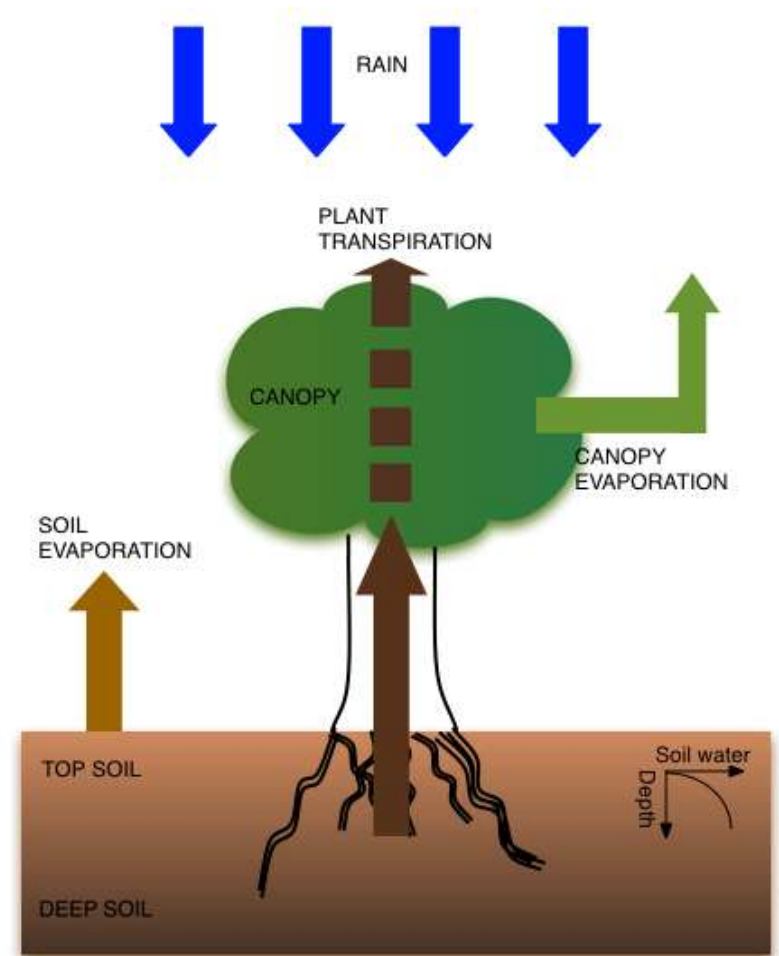


# Canopy capacity in JULES

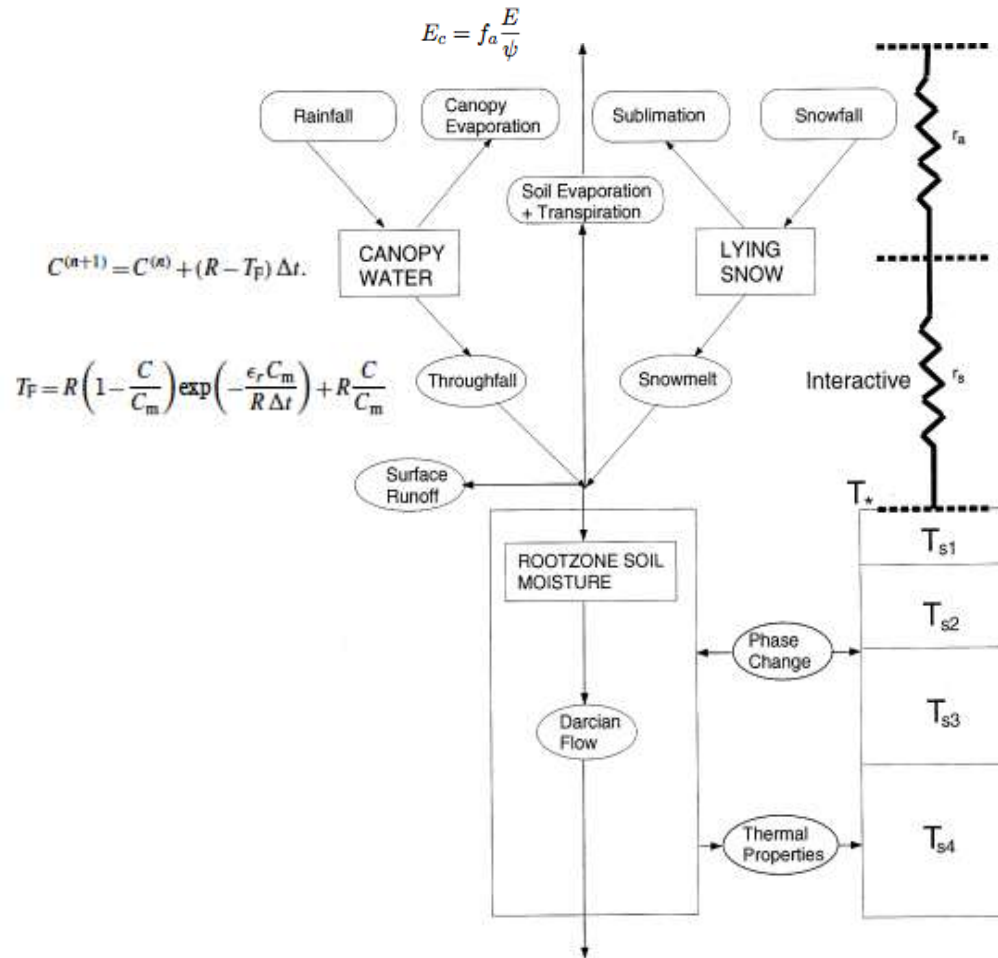
Emily Black, Marie-Estelle Demory,  
Pier Luigi Vidale, Anne Verhoef and  
Catherine Van den Hoof

# Outline

- Background and motivation
- Effect of different formulations of canopy capacity on latent heat simulation in JULES
- The knock-on effect on soil moisture and NPP
- Conclusions and recommendations



# The hydrological cycle in JULES

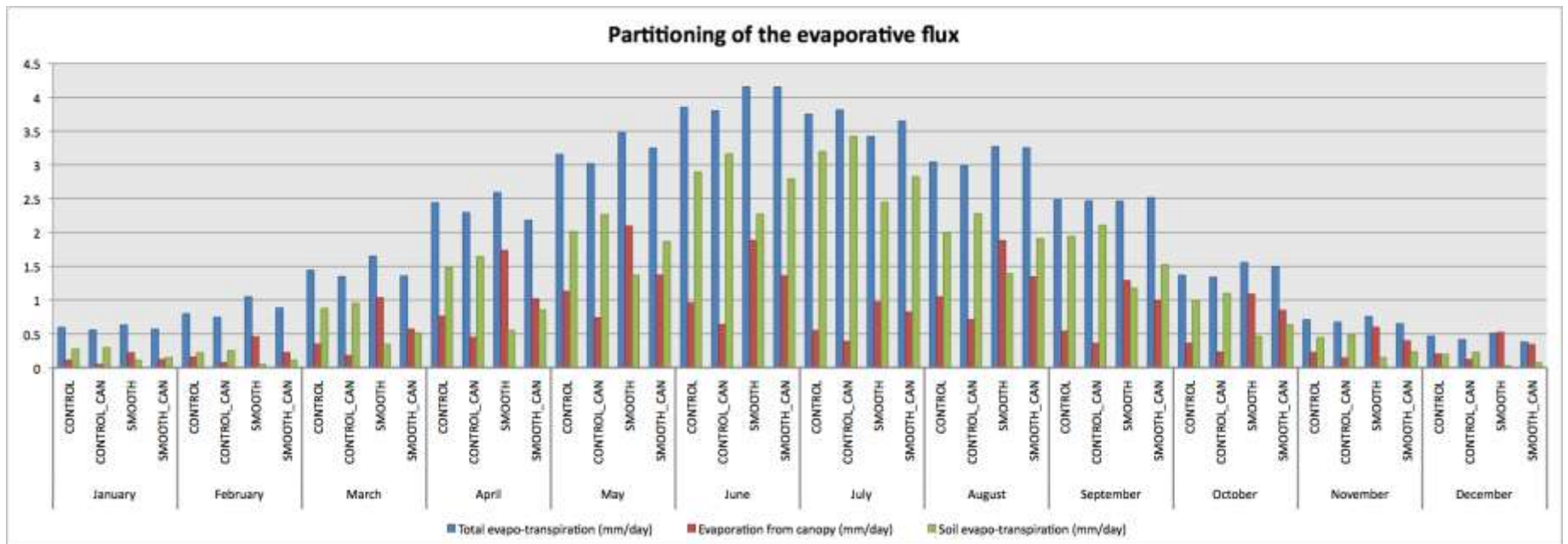
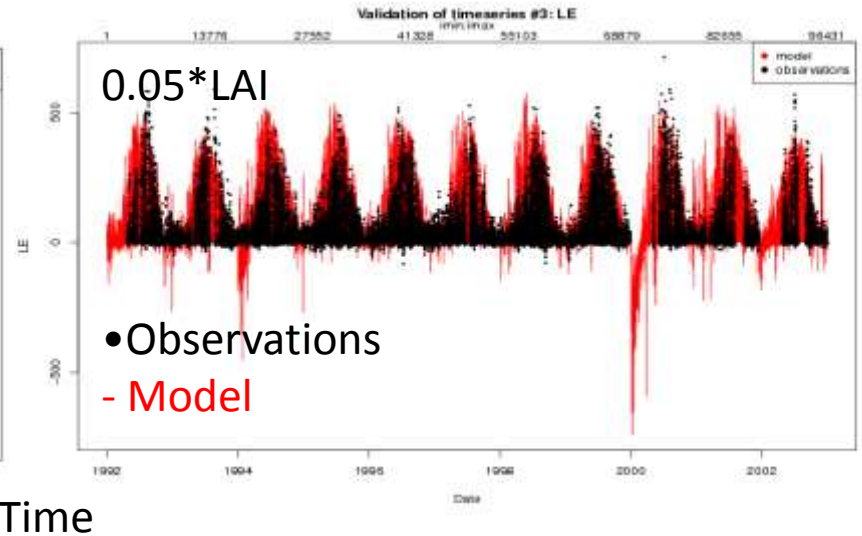
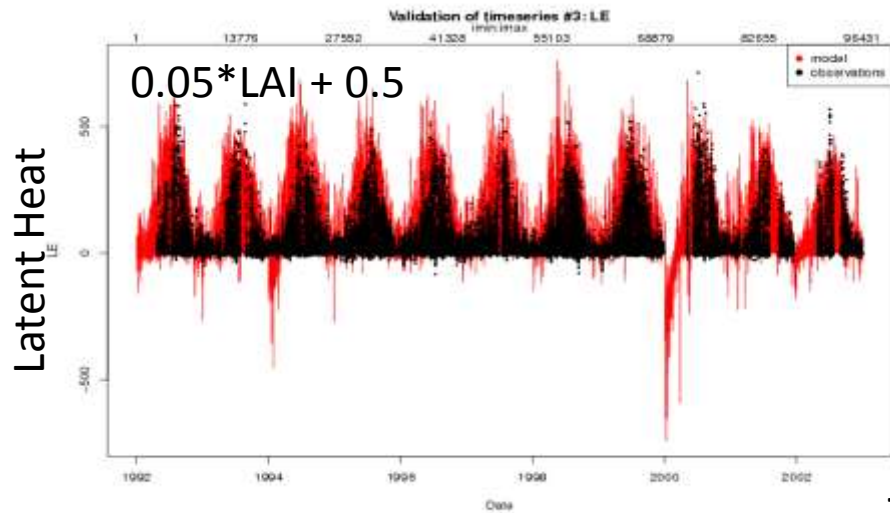


It is important to get canopy capacity right because it determines how moisture is exchanged between surface and atmosphere within a time step, and hence the degree of land-surface - atmosphere coupling.

Default canopy capacity in JULES is  $0.5 + 0.05 * LAI$ . BUT:

1. In temperate regions, the seasonal variation in canopy capacity may not be well-represented
2. The canopy capacity may be too high.
3. It is at odds with other LSMs (eg SiB and NCAR-LSM), which define canopy capacity as  $A * LAI$

# Canopy capacity and evaporative fluxes



# Methodology

## Driving/validation data:

- Benchmarking fluxnet sites
- Hourly/half hourly meteorological driving data
- Vegetation/soils based on observations
- Short data periods (up to 13 years)

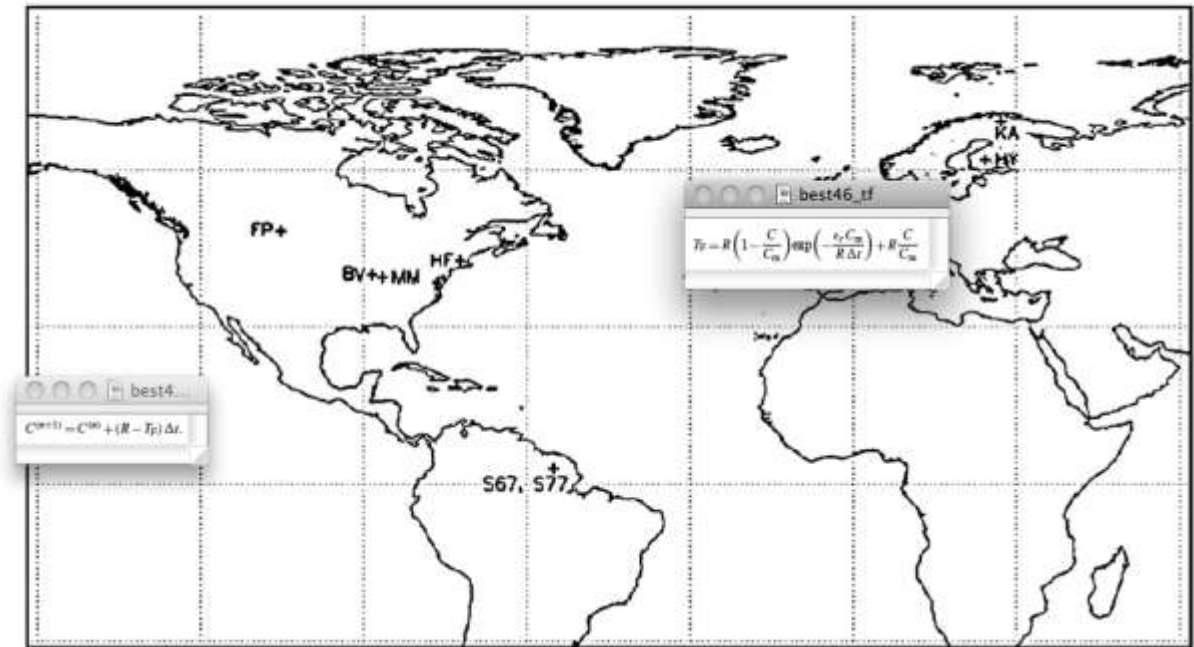
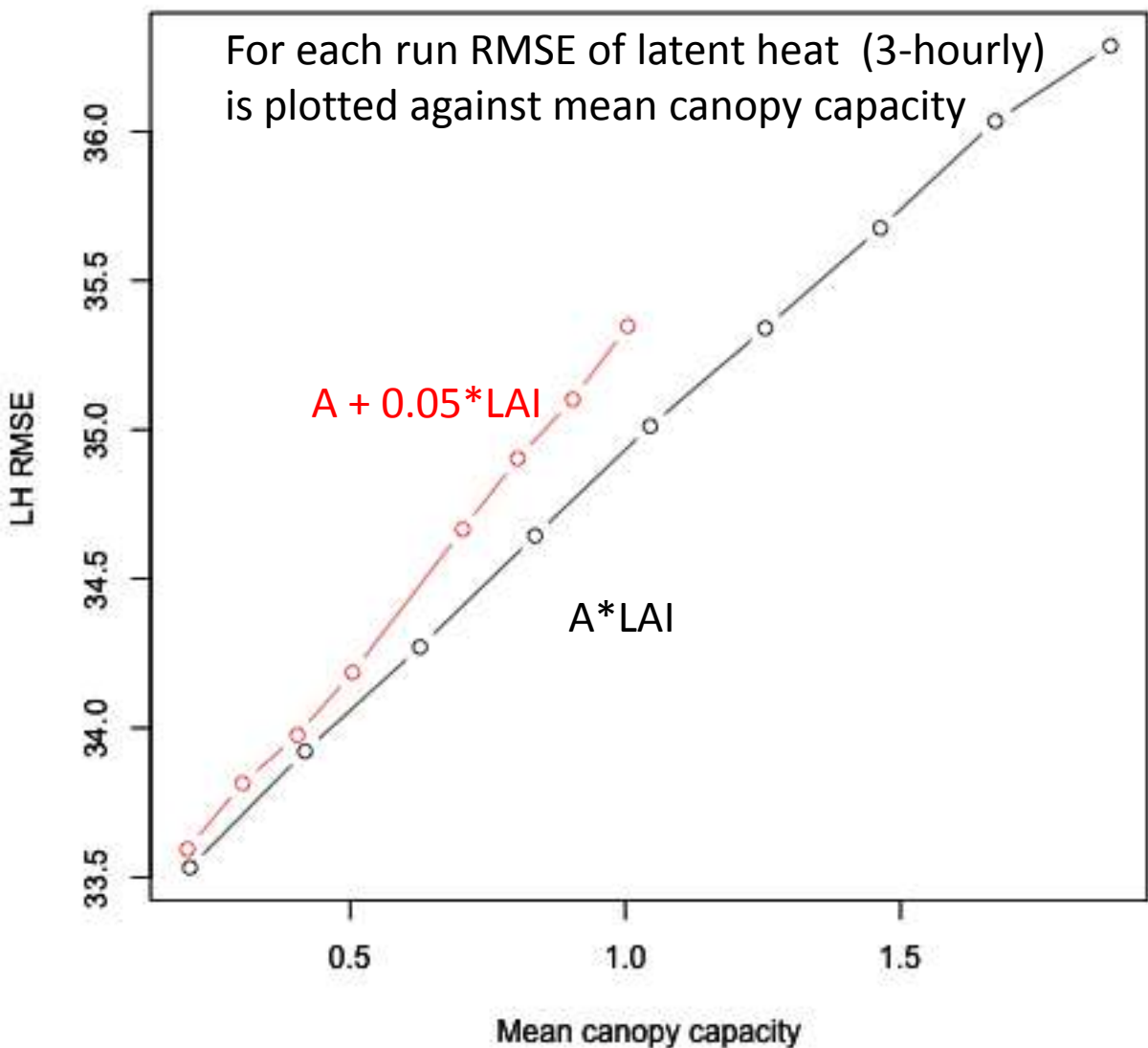


FIG. 1| Location of the FLUXNET sites used in this study labeled as follows: FP = Fort Peck; BV = Bondville; MM = Morgan Monroe; HF = Harvard Forest; S67 = Santarem km 67, Brazil; S77 = Santarem km 77, Brazil; KA = Kaamenen; HY = Hyytiälä; TH = Tharandt; and ES = El Saler.

## JULES set up:

- Phenology ON
- Canopy model 4
- Full spin up

# Effect of canopy capacity function on the simulation of Latent Heat

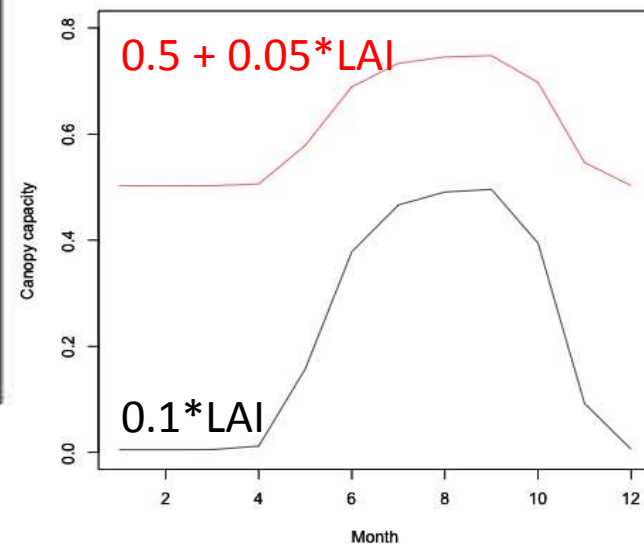


JULES forced with fluxnet data. Canopy capacity function of the form:

(i)  $A + 0.05 * LAI$

(ii)  $A * LAI$

The constant 'A' varies from 0.1-0.9



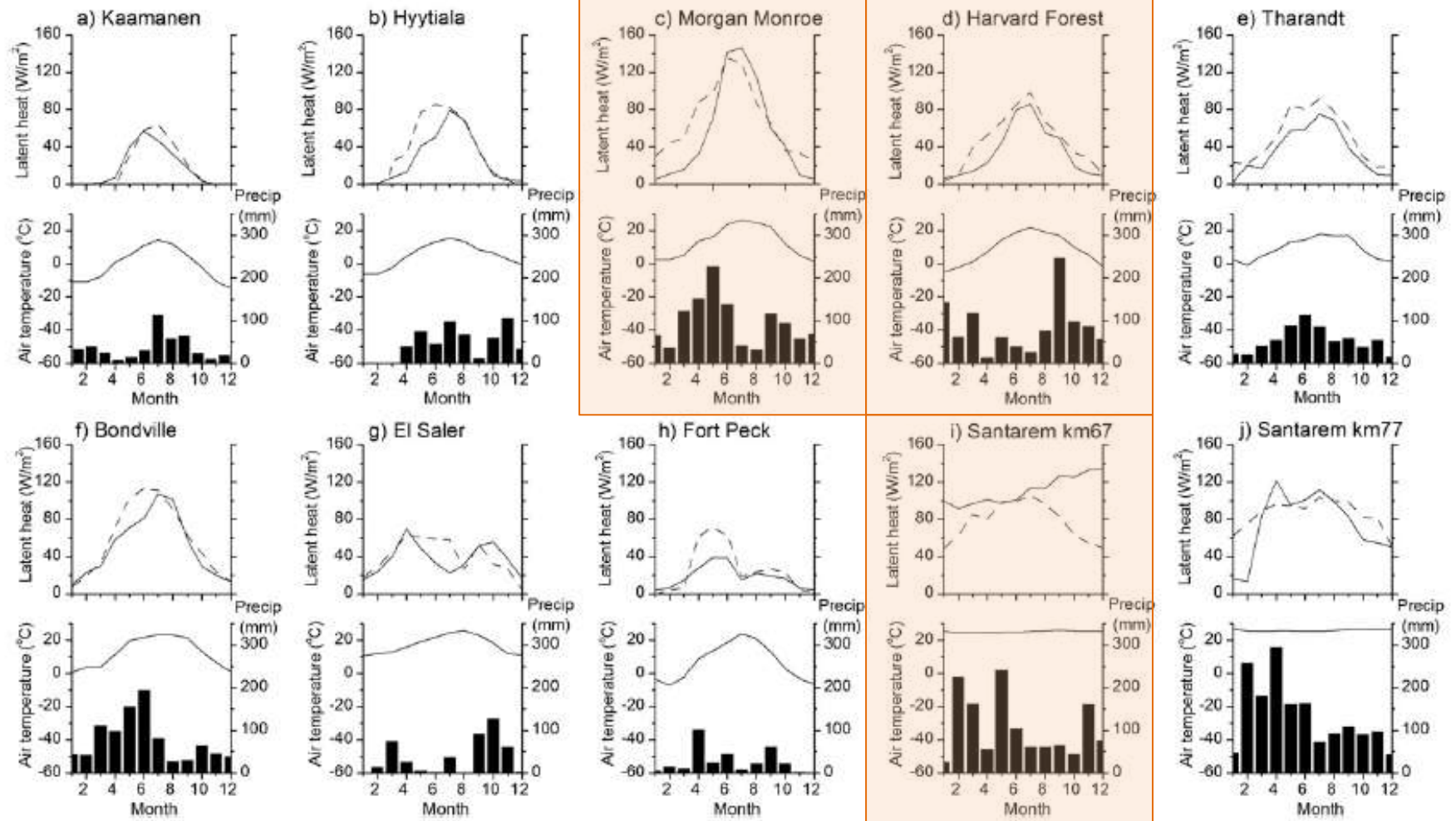
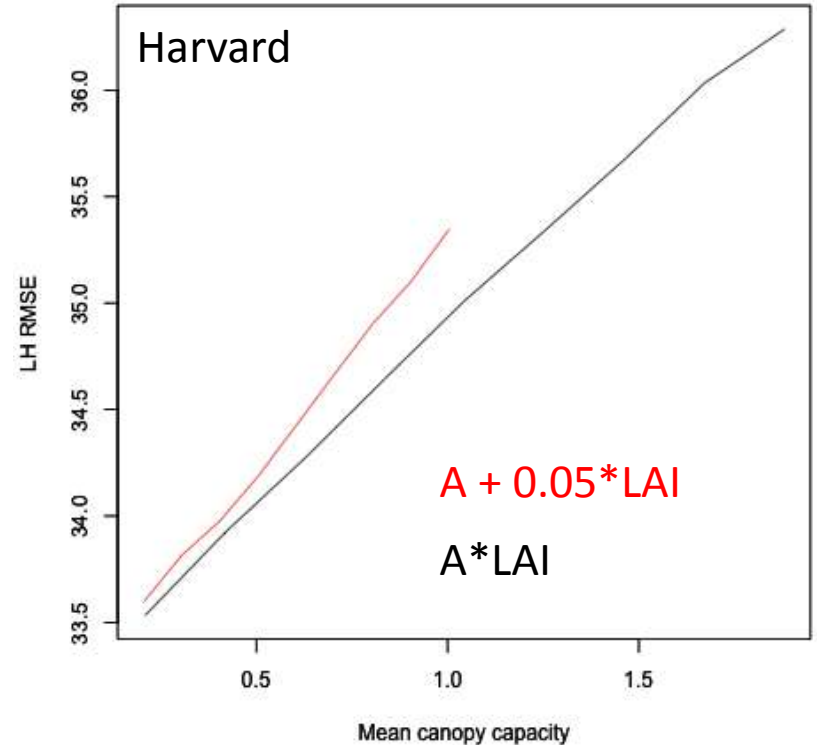
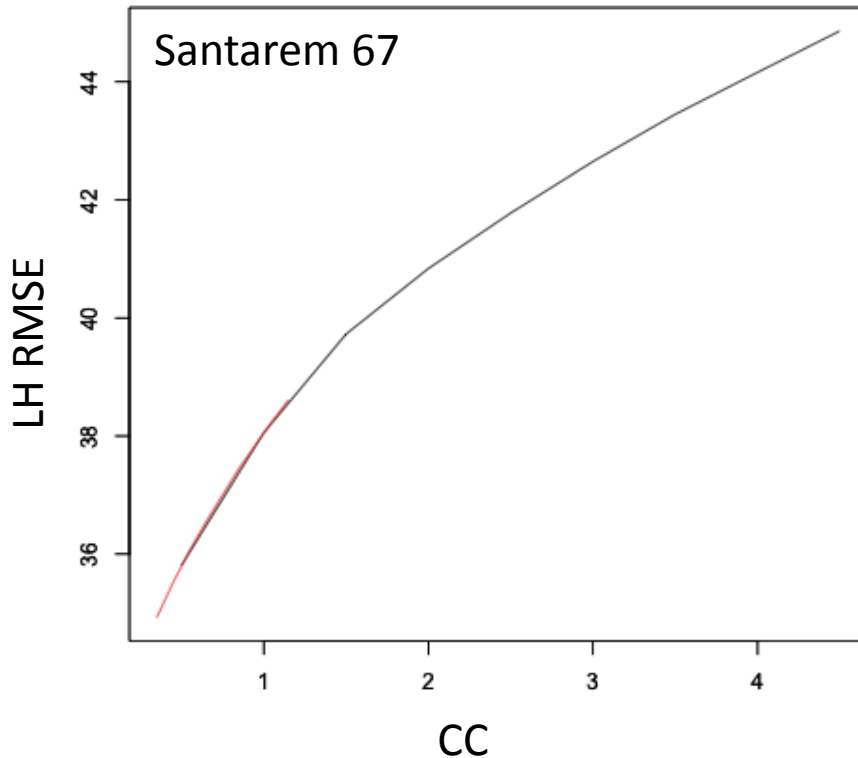


FIG. 2. Monthly-mean observed evaporation ( $E_{OB}$ , solid thick line) and modeled evaporation ( $E_{MOD}$ , solid thick broken line) air temperature (thin solid line), and rainfall (bar chart) for the following FLUXNET sites: (a) Kaamanen, (b) Hyytiälä, (c) Morgan Monroe, (d) Harvard Forest, (e) Tharandt, (f) Bondville, (g) El Saler, (h) Fort Peck, (i) Santarem km 67, and (j) Santarem km 77.

BT

# Broad-leaf tree sites



- Most error arises from the systematically high latent heat in JULES – not from canopy capacity.
- The higher the canopy capacity, the higher the error
- It is evident, however that the errors are lower for canopy capacity of the form  $A * LAI$



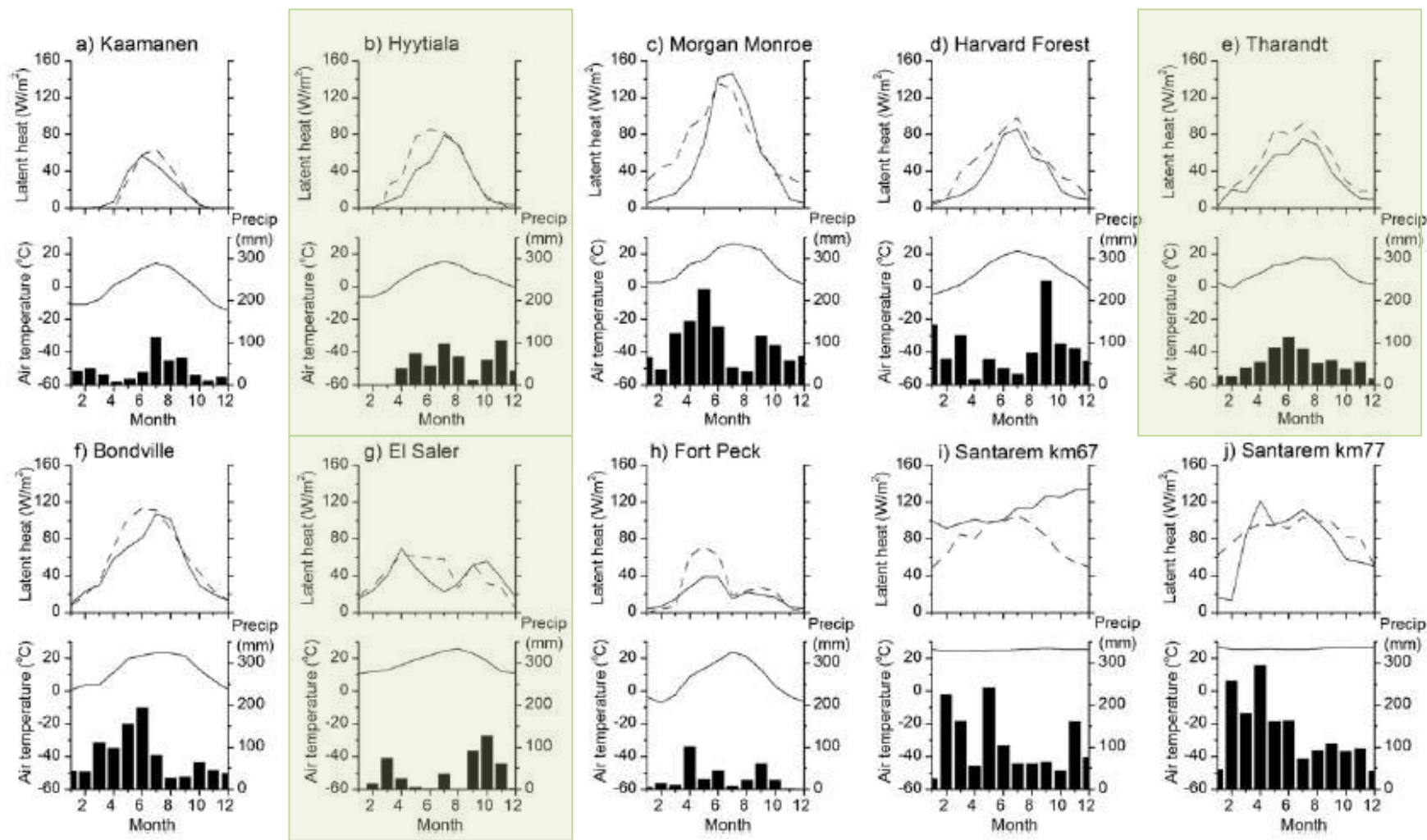
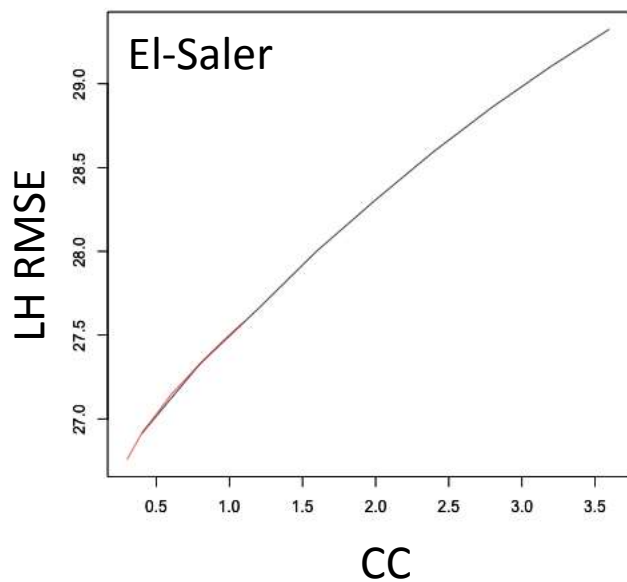
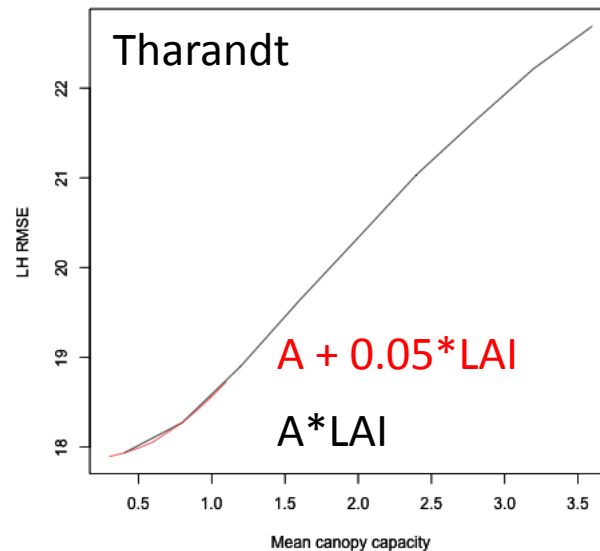
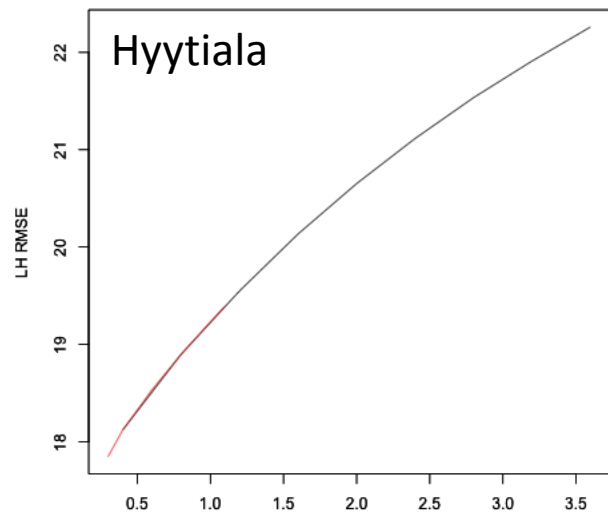


FIG. 2. Monthly-mean observed evaporation ( $E_{OB}$ , solid thick line) and modeled evaporation ( $E_{MOD}$ , solid thick broken line) air temperature (thin solid line), and rainfall (bar chart) for the following FLUXNET sites: (a) Kaamanen, (b) Hyytiälä, (c) Morgan Monroe, (d) Harvard Forest, (e) Tharandt, (f) Bondville, (g) El Saler, (h) Fort Peck, (i) Santarem km 67, and (j) Santarem km 77.

NT

# Needle-leaf tree sites



Not surprisingly, the form of the canopy capacity function does not matter for needle-leaf trees (LAI does not vary seasonally in JULES)

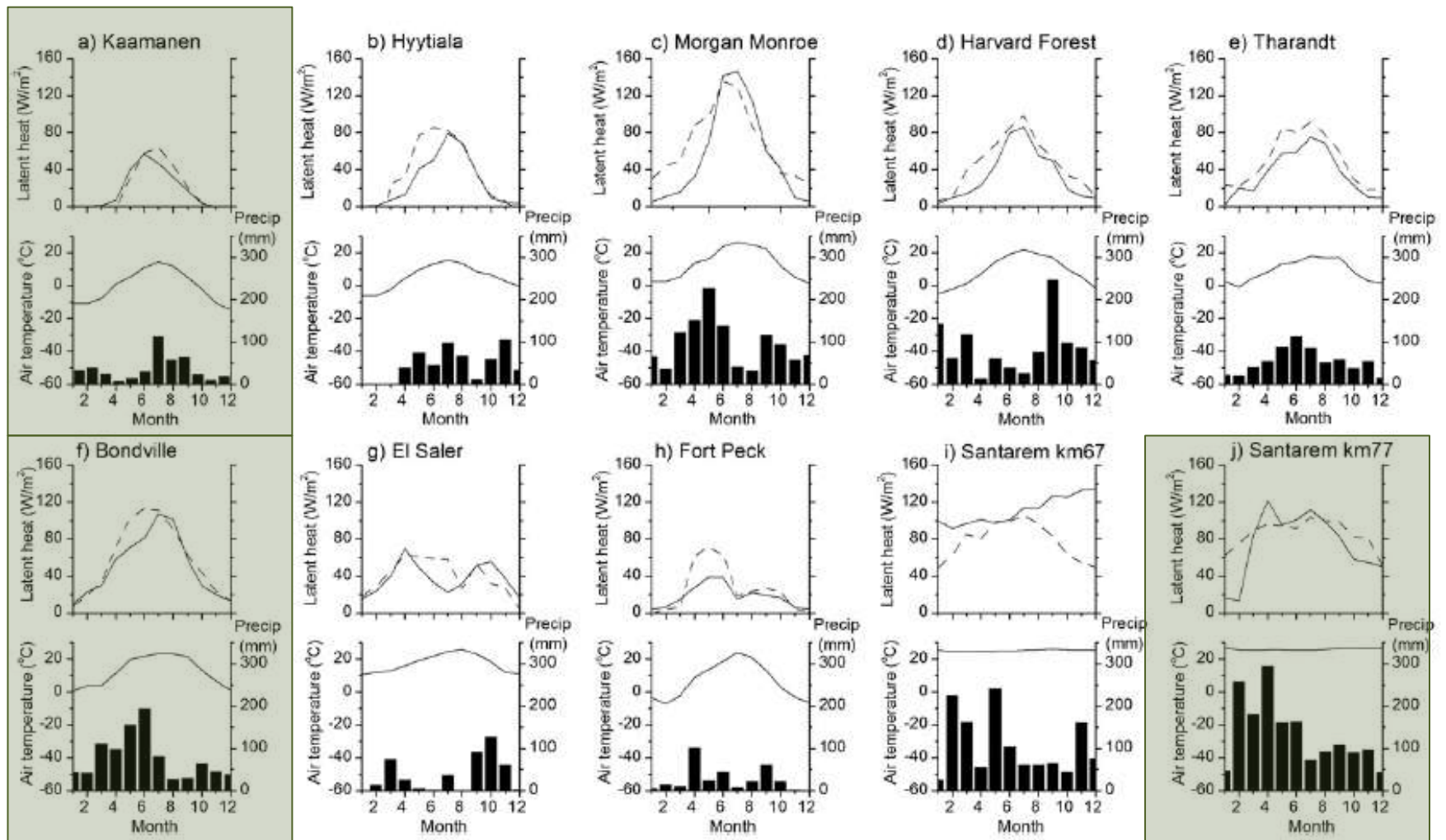
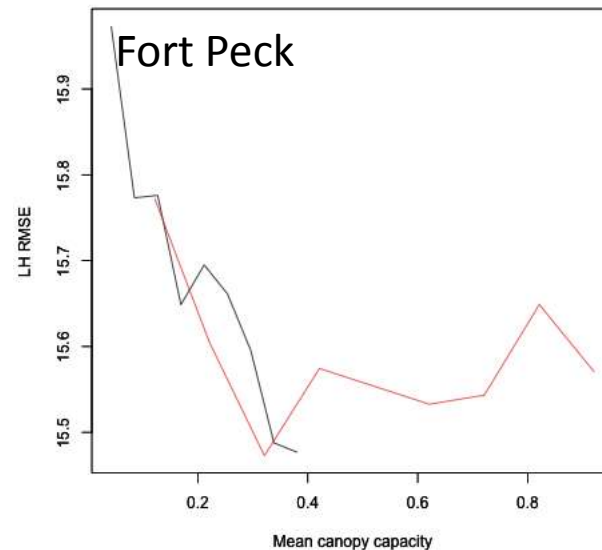
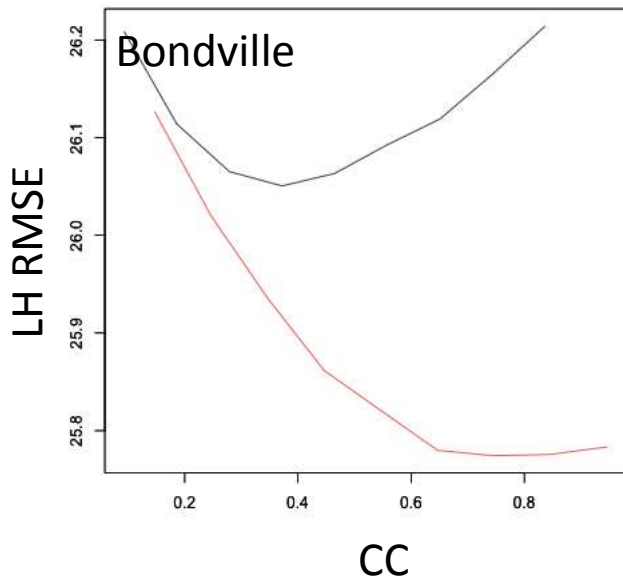
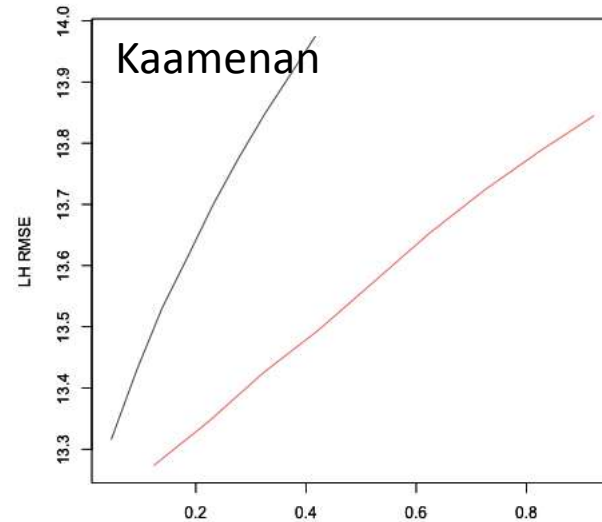
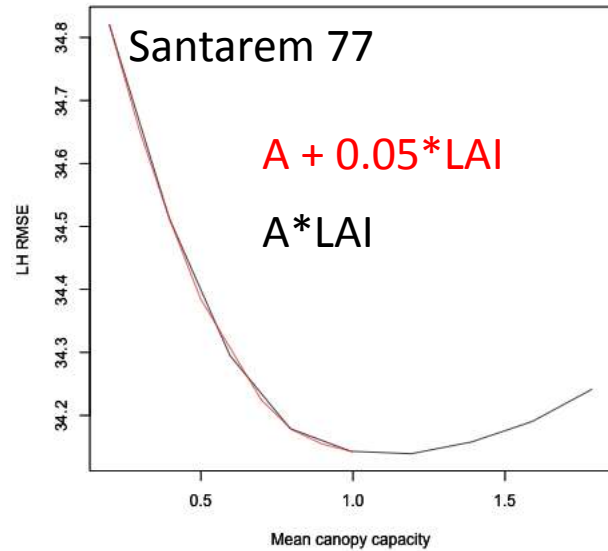


FIG. 2. Monthly-mean observed evaporation ( $E_{OB}$ , solid thick line) and modeled evaporation ( $E_{MOD}$ , solid thick broken line) air temperature (thin solid line), and rainfall (bar chart) for the following FLUXNET sites: (a) Kaamanen, (b) Hyytiälä, (c) Morgan Monroe, (d) Harvard Forest, (e) Tharandt, (f) Bondville, (g) El Saler, (h) Fort Peck, (i) Santarem km 67, and (j) Santarem km 77.

# C3 Grass sites

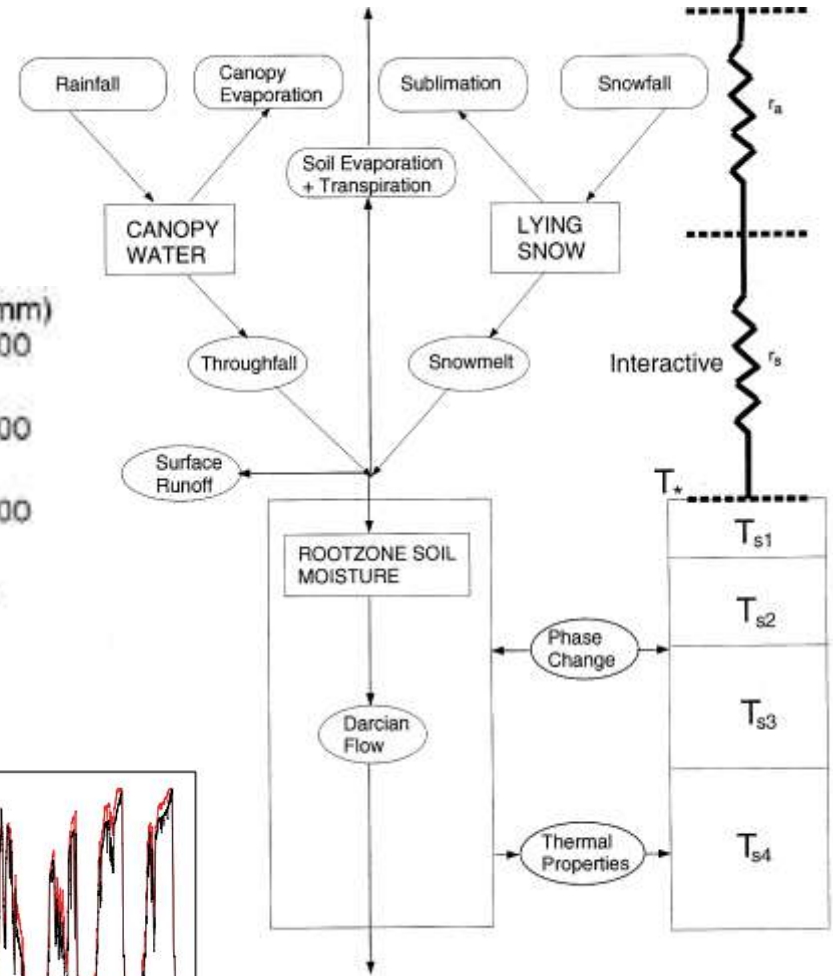
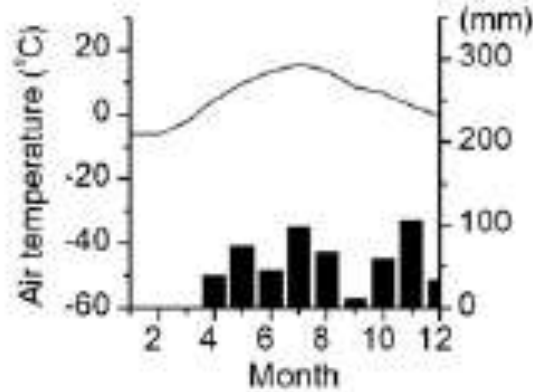
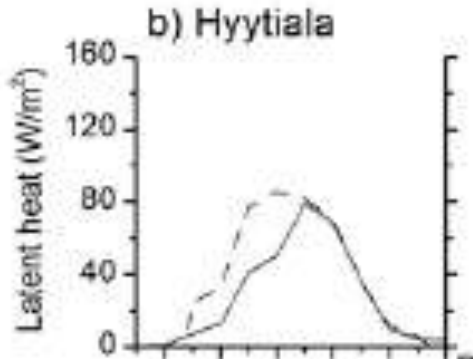


C3 grass results are difficult to interpret. There is considerable **variability between sites and regions.**

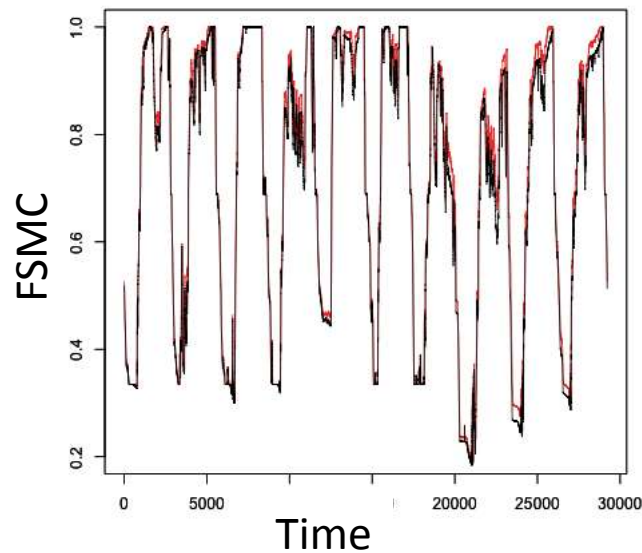
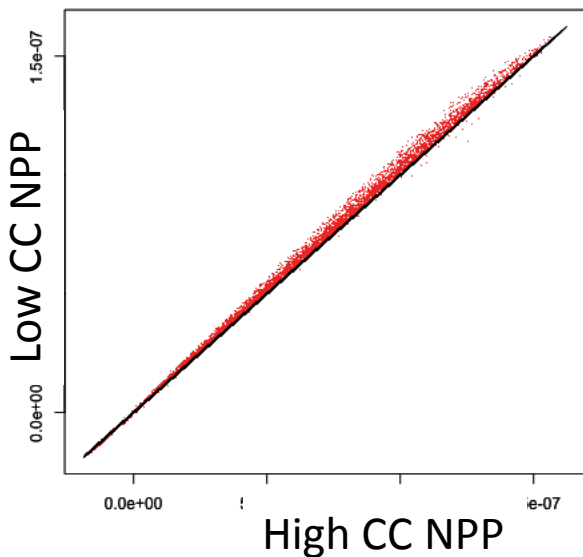
This may be related to JULES' simulation of the seasonal cycle in temperate regions.

# Vegetation and CC

Hyytiala: Needle leaf trees

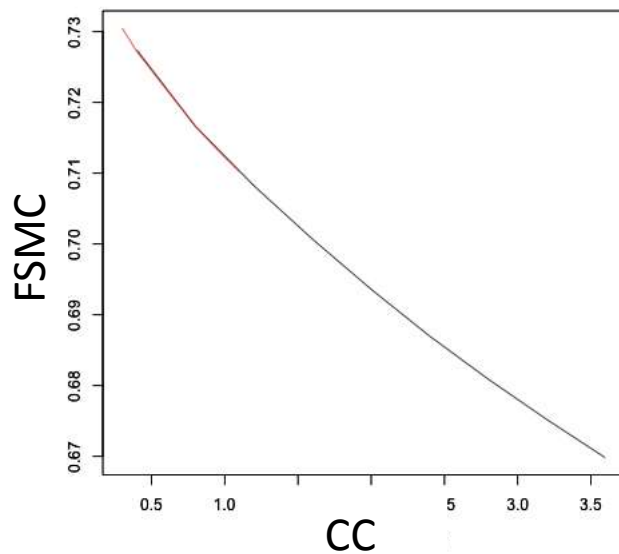
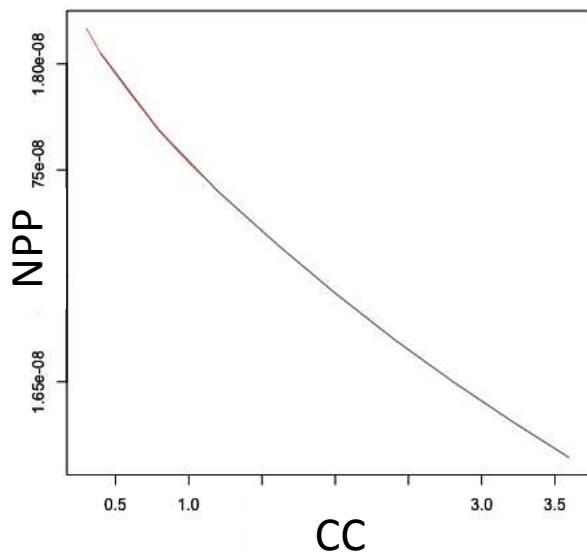
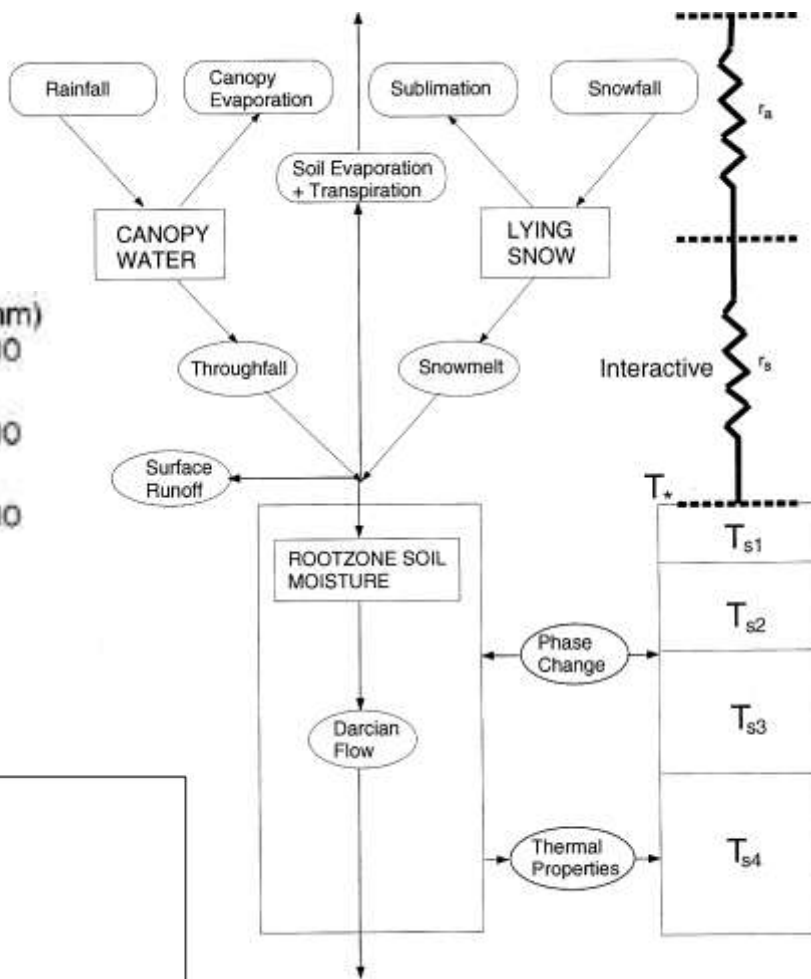
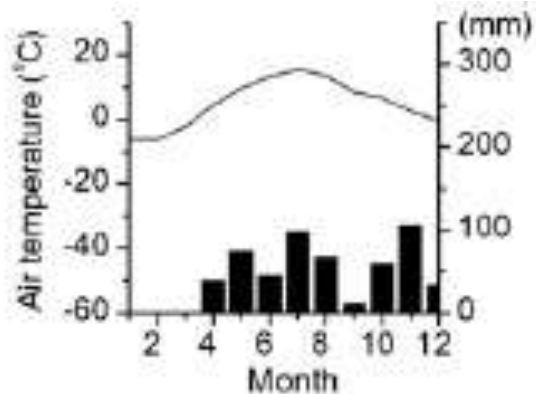
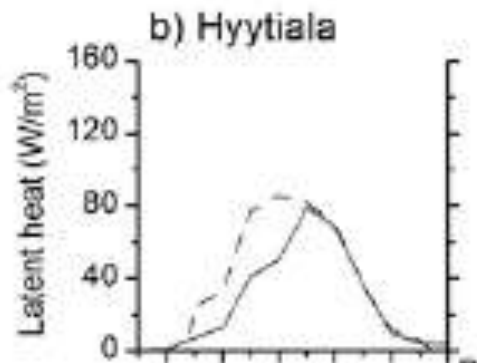


-High Canopy Capacity  
-Low Canopy Capacity



# Vegetation and CC

Hyttiala: Needle leaf trees



# Conclusions

- It is important to get canopy capacity right, because of its importance for land-atmosphere coupling.
- However, in the cases of needle-leaf and broadleaf trees, **canopy capacity is not the primary cause of error in latent heat**. Generally, reducing the canopy capacity will reduce errors mainly by compensating for excessively high evapo-transpiration.
- It is evident that canopy capacity of the form  $A \cdot LAI$  is more appropriate in temperate BT/NT regions because it introduces a more realistic seasonal cycle in canopy capacity, which improves the seasonal cycle in latent heat.
- In the case of grasses, the situation is more complex. Canopy capacity has an effect on error over and above its impact on reducing evapo-transpiration.

# Recommendations

- For broadleaf and needle leaf trees, canopy capacity should be defined as  $A * LAI$ , where  $A$  is between 0.1 and 0.2. This can be done in the .jin file
- Further work should be carried out to determine the best parameterization for grasses. For now, it would be best to stick to  $0.5 + 0.05 * LAI$
- Sorting out the canopy capacity will not solve JULES' problem of high evapo-transpiration. Further work on the hydrological cycle in JULES (eg during SWELTER) is required.